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a model-independent technique has been discussed to interpret self-potential anomalies over a pole, a horizontal line of poles, a horizontal cylinder, a sphere and an inclined sheet. the present method involves computation of second order derivatives with respect to cartesian coordinate system. an equation, which is independent of the source geometry, has been used to compute the unknown source parameters. a procedure has been discussed to suppress the high frequency noises present in the observations to achieve derivatives of the signal with reasonably correct amplitude and low noise content. a study on the reliability and the applicability of the technique reveals that for accurate determination of depth and structural index (si) of a source, the ratio of depth to data spacing should be 5 and above. further, si computation from kx is more reliable than that by kz. a random error of 10% in simulated anomalies has no significant effect on the computed model parameters except for a horizontal line of poles. the applicability of the proposed technique has been demonstrated through four field examples, with different levels of complexities, adopted from published literature. the aim of this study is to determine the k-excursion ratio to accurately determine the si of the geologic structure. the k-excursion ratio is defined as the ratio of the source depth to the data spacing. the k-excursion ratio is determined by two procedures; (a) first, the si is calculated from kx and then the k-excursion ratio is determined from the ratio of kx and depth, and (b) first, the si is calculated from kz and then the k-excursion ratio is determined from the ratio of kz and depth.

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seismic velocity data in both the transition zone and the iberia margin show that the velocity structure is characterized by distinct layers of high and low velocity (fig. f6). at the start of the transition zone, the high velocity layer is ~50 km thick, and is bounded on the west by a sharp gradient to a low-velocity zone that extends 60 km southwestward. this low-velocity zone is bounded on the east by a steep gradient to the high-velocity layer. on the iberia margin, the high-velocity layer appears to be thicker (100 km) and to extend further west than on the transition zone. this is likely a result of the weakness of the low-velocity zone in the north-central transition zone, which is bounded on the west by a steep low-velocity zone that extends 100 km southeastward (fig. f7b). a low-velocity zone that is >80 km wide separates these two high-velocity layers. the transition zone does not extend as far south as the iberia margin, where the high-velocity layer is ~70 km thick and does not extend as far west as in the central segment (figs. f3b, f7a). these velocity trends agree with those shown by the sp anomalies in the transition zone and iberia margin, suggesting that the transition-zone crust is older than the iberia margin (fig. f4). figure f4 shows the sp anomalies with superimposed low-velocity zones from transect 2. i have adjusted the transect to show only the central portion of the transition zone, where a low-velocity zone is the most prominent feature. this figure shows that the transition zone is characterized by a series of low-velocity zones that are separated by high-velocity zones. at the start of the transition zone (fig. f4a), the low-velocity zone is ~50 km wide, and it extends 60 km southwestward. the eastern and western margins of the low-velocity zone are marked by steep low-velocity gradients.

these gradients are thought to represent subducting slabs, and the high-velocity zone at the western edge of the transition zone has been interpreted to be a down-dip continuation of the iberia margin (srivastava et al., 2000). a low-velocity zone, which is thought to be connected to the iberia margin, separates the two high-velocity zones in the central segment. this low-velocity zone is ~80 km wide, and it also separates the low-velocity zones in the transition zone from the high-velocity zones in the iberia margin. 5ec8ef588b

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